

An ocean-ice model study of the post-1995 collapse of the North Atlantic subpolar gyre

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Outline

- Ocean model and its configuration.
- Atmospheric forcing.
- Spinup procedure.
- Investigation of the post-1995 collapse of the sub polar gyre.



Ocean model and its configuration

The ocean model is MICOM with several modifications:

- Incremental remapping used for advection of layer thickness and tracers.
- Both temperature and salinity are prognostic variables.
- No linearization of the bottom pressure.
- New formulation of the pressure gradient force.
- New equation of state.
- Modified barotropic/baroclinic mode splitting.
- Parameterizations of shear instability and gravity current diapycnal mixing are included.
- Improved conservation.

Configuration

- Global domain, non-eddy resolving. Some experiments done with enhanced resolution in the North Atlantic/Arctic region.
- 35 isopycnic layers.
- Reference pressure at 2000 db.



Atmospheric forcing

- Daily forcing fields from either NCEP/NCAR or ERA40 reanalysis.
- The reanalysis air-sea fluxes of momentum, heat, and freshwater are reproduced exactly if the models ocean state is identical to the ocean state of the reanalysis.
- When the ocean state differs, the turbulent fluxes are modified consistent with the Fairall et al. (1996) bulk parameterization and long-wave radiation is modified consistent with the Berliand and Berliand (1952) parameterization.
- The method is described in Bentsen and Drange (2000).

$$\tau^d = -\rho_a C_D^d (u_a^d)^2$$

$$Q_H^d = \rho_a c_p C_H^d u_a^d (T_s^d - T_a^d)$$

$$Q_L^d = \rho_a L_e C_E^d u_a^d (q_s^d - q_a^d)$$

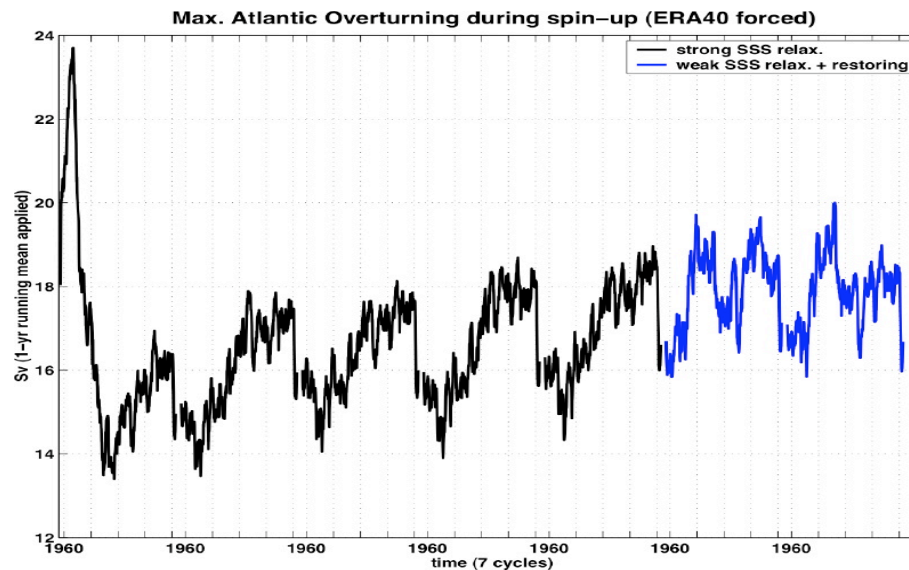
$$\tau^m = -\rho_a C_D^m (u_a^d)^2$$

$$Q_H^m = \rho_a c_p C_H^m u_a^d (T_s^m - T_a^d)$$

$$Q_L^m = \rho_a L_e C_E^m u_a^d (q_s^m - q_a^d)$$

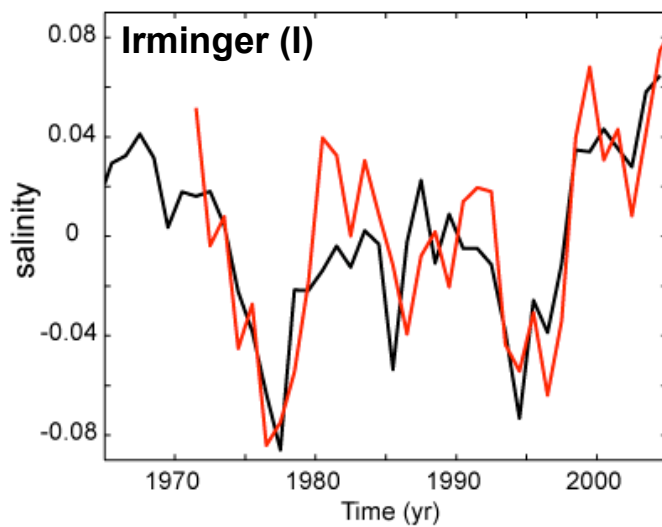
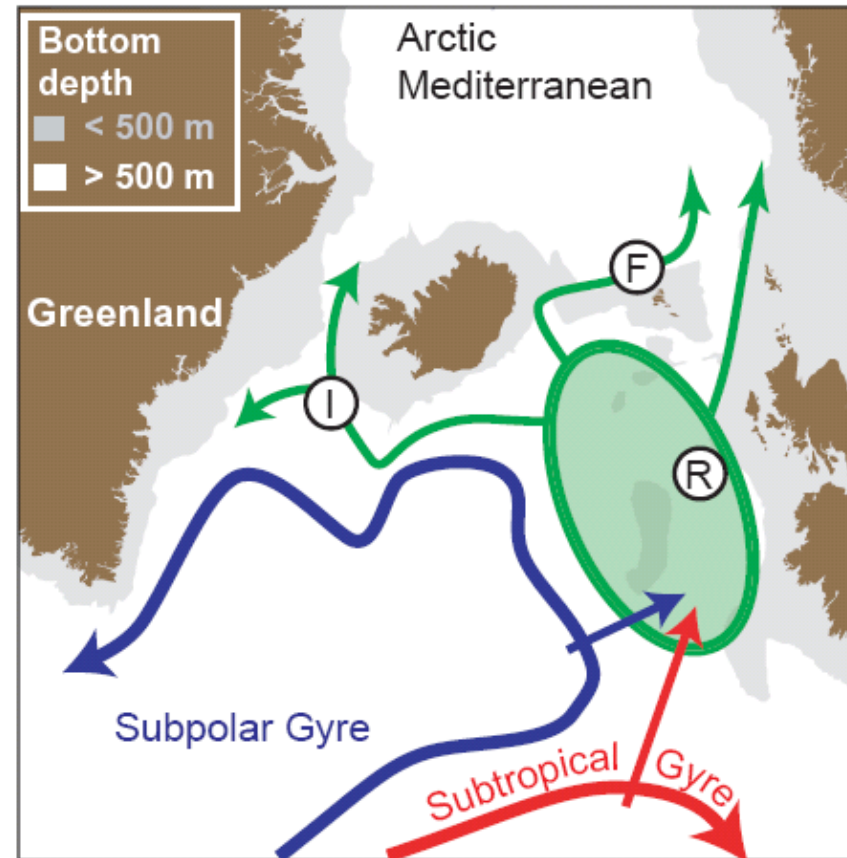
Spinup procedure

- Minimum full 4 reanalysis cycles with daily forcing; usually 6 full cycles (~300 yr).
- Starts with strong SSS-relaxation; 30 days e-folding time scale for 50 m thick mixed layer; limited to $|\Delta\text{SSS}<0.5|$ everywhere; no relaxation under sea-ice.
- Diagnose SSS-nudging when model is steady (5th or 6th cycle); applying diagnosed SSS-fluxes for the production runs with weak Newtonian relaxation (360 days time scale for 50 m ML and $|\Delta\text{SSS}<0.5|$).
- Temperature relaxation is not used.

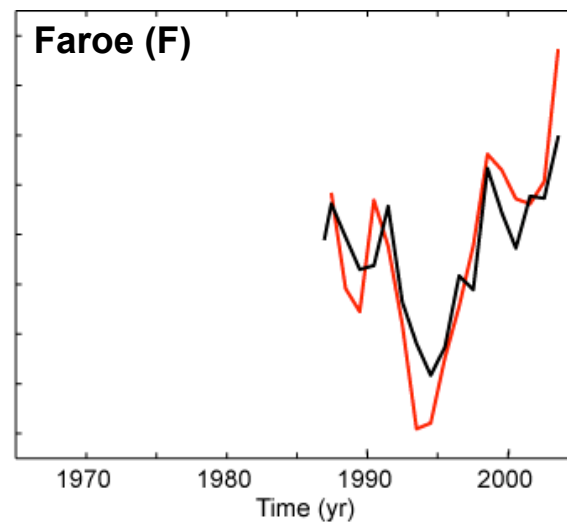
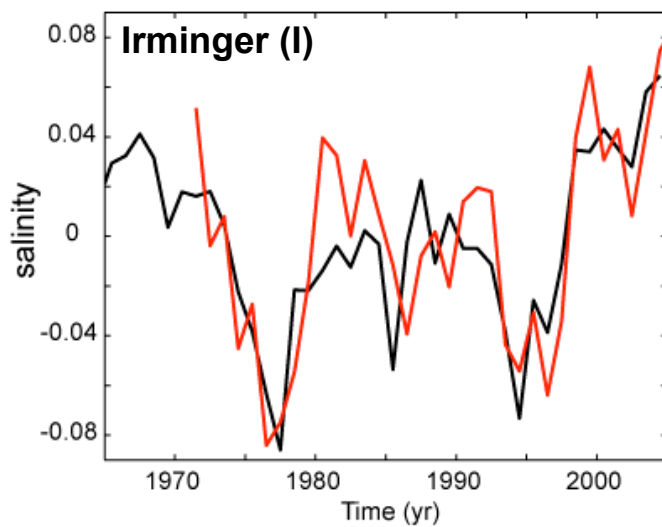
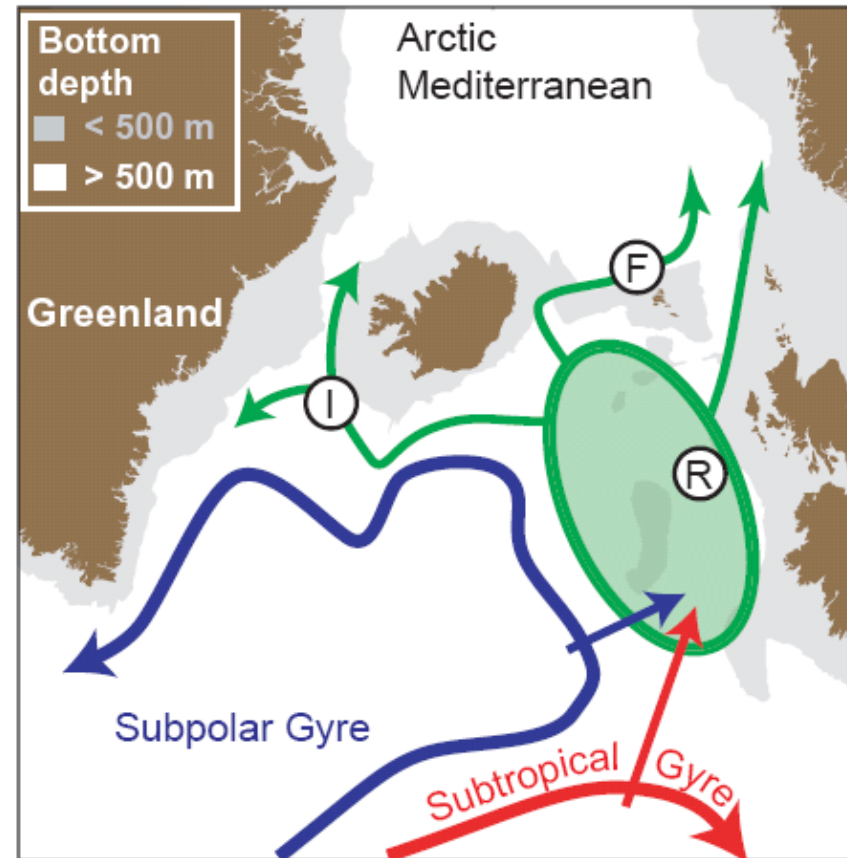


Hatun et al., *Science* (2005); *Prog. Ocenogr.* (2009)

Observed and **simulated** salinity anomalies at three locations in the northern North Atlantic



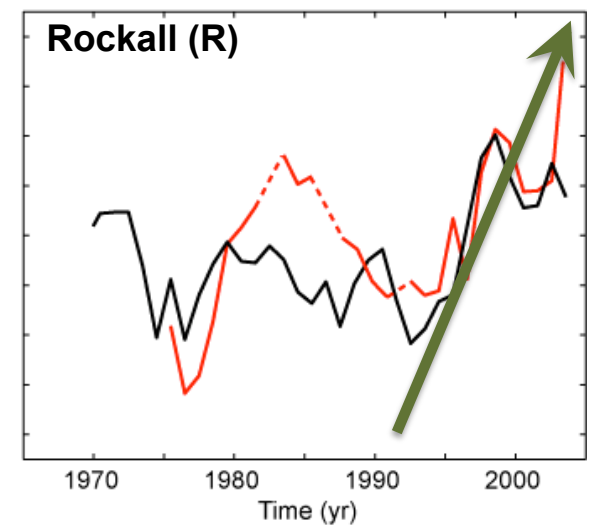
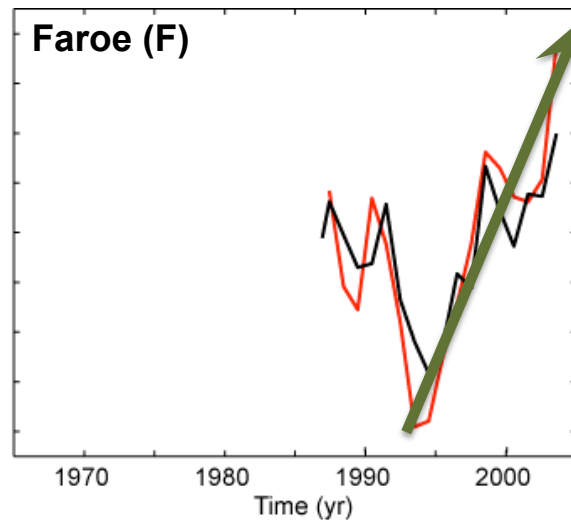
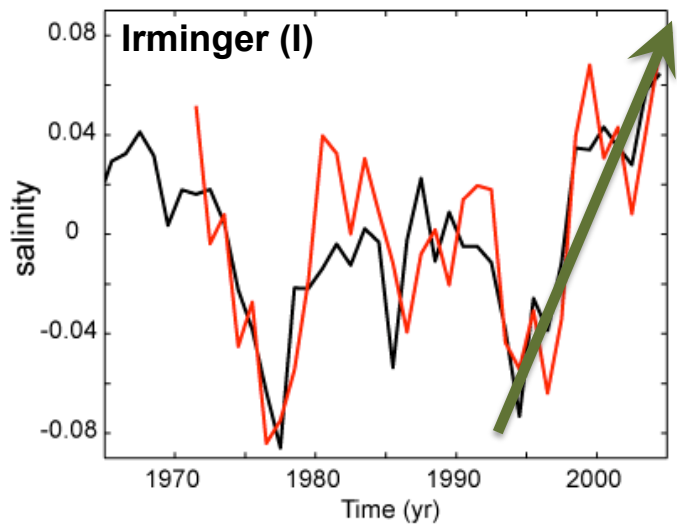
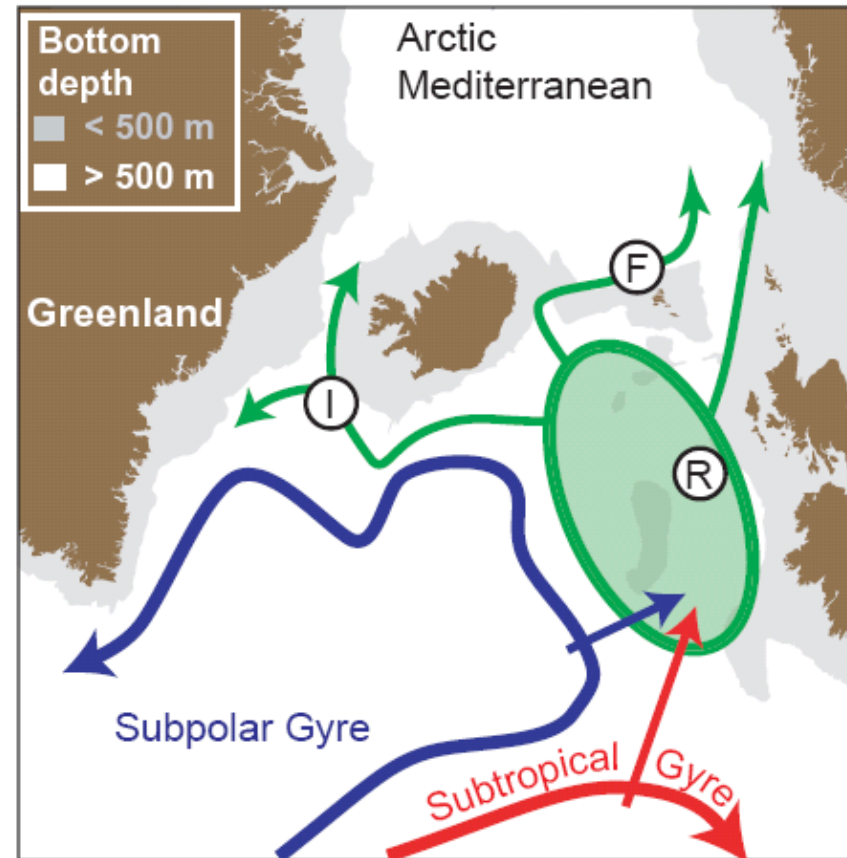
Observed and **simulated** salinity anomalies at three locations in the northern North Atlantic



Observed and **simulated** salinity anomalies at three locations in the northern North Atlantic

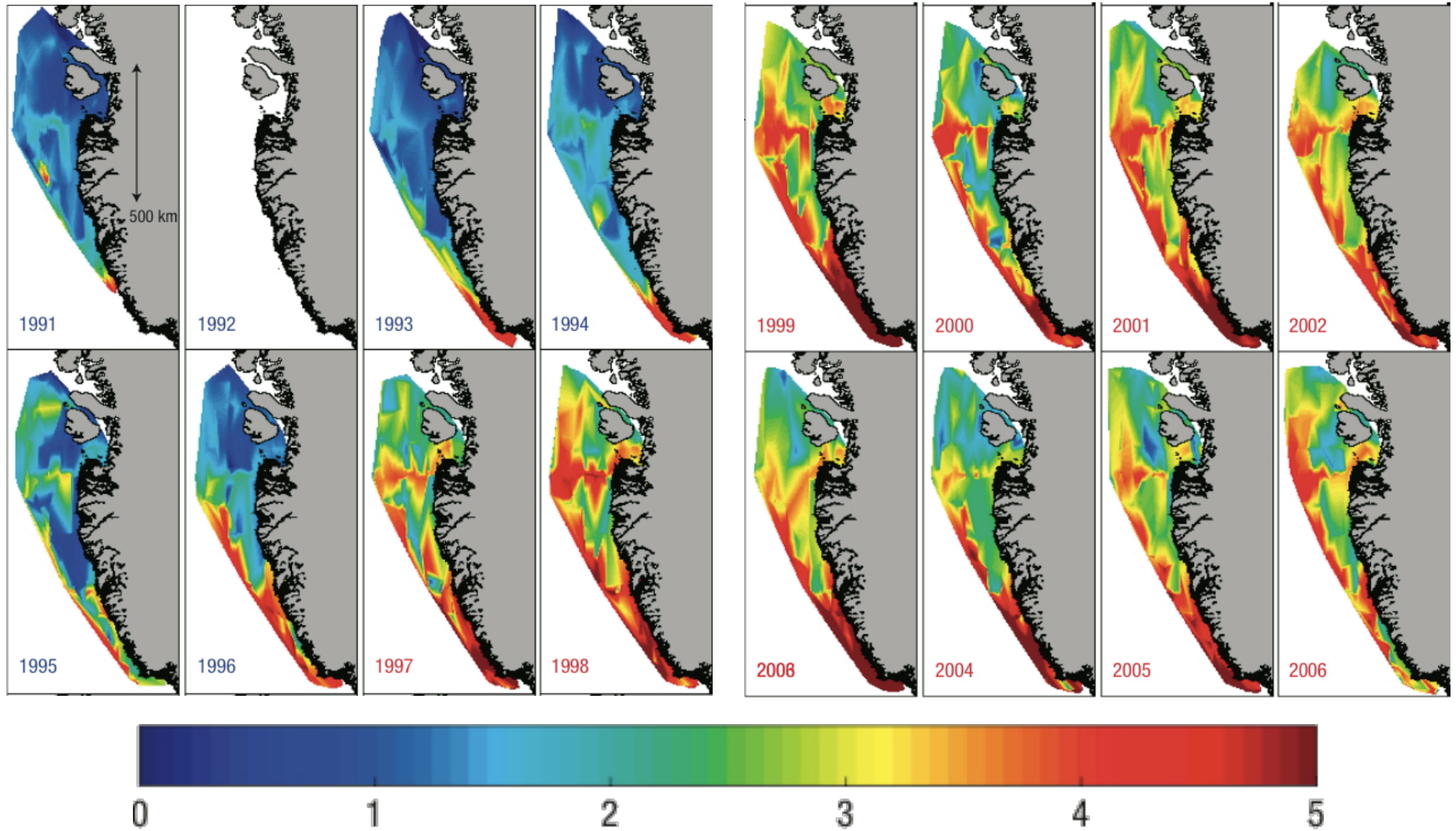
Why post-95 change?

(T, S, marine biota)



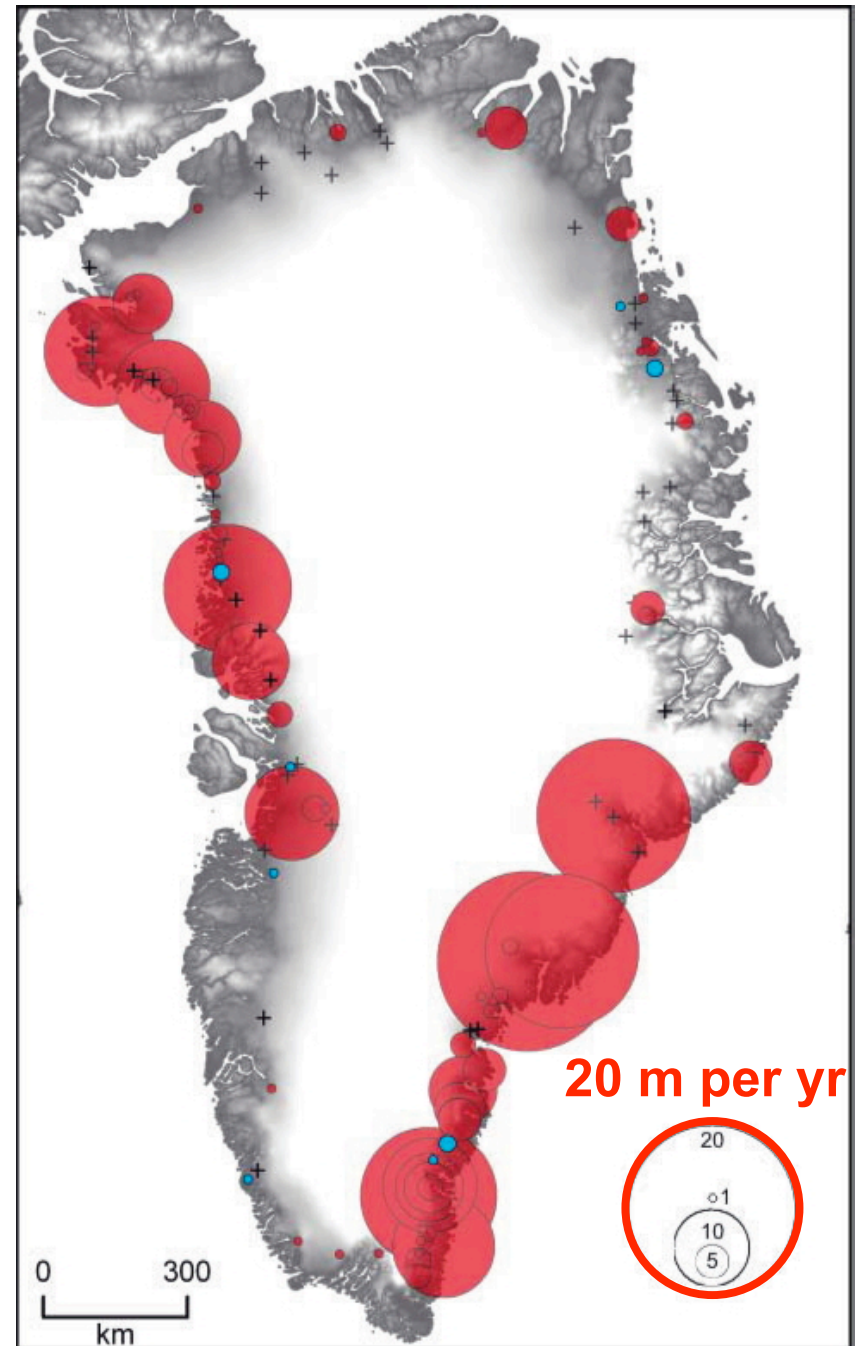
Associated warming off the coast of Greenland

(150-600 m)

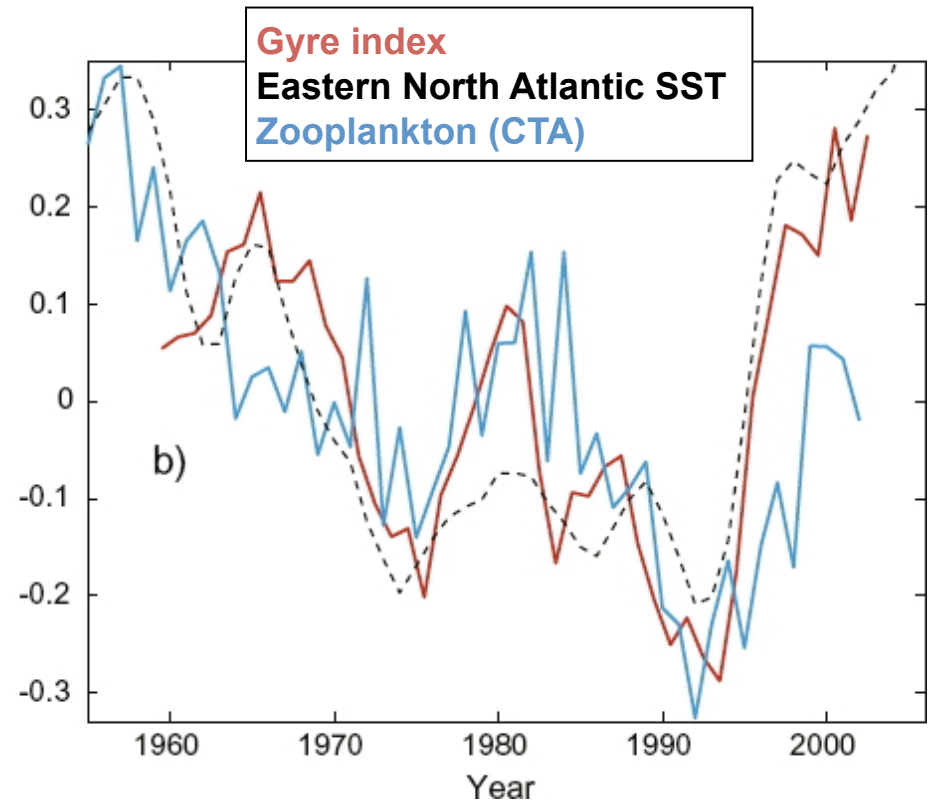
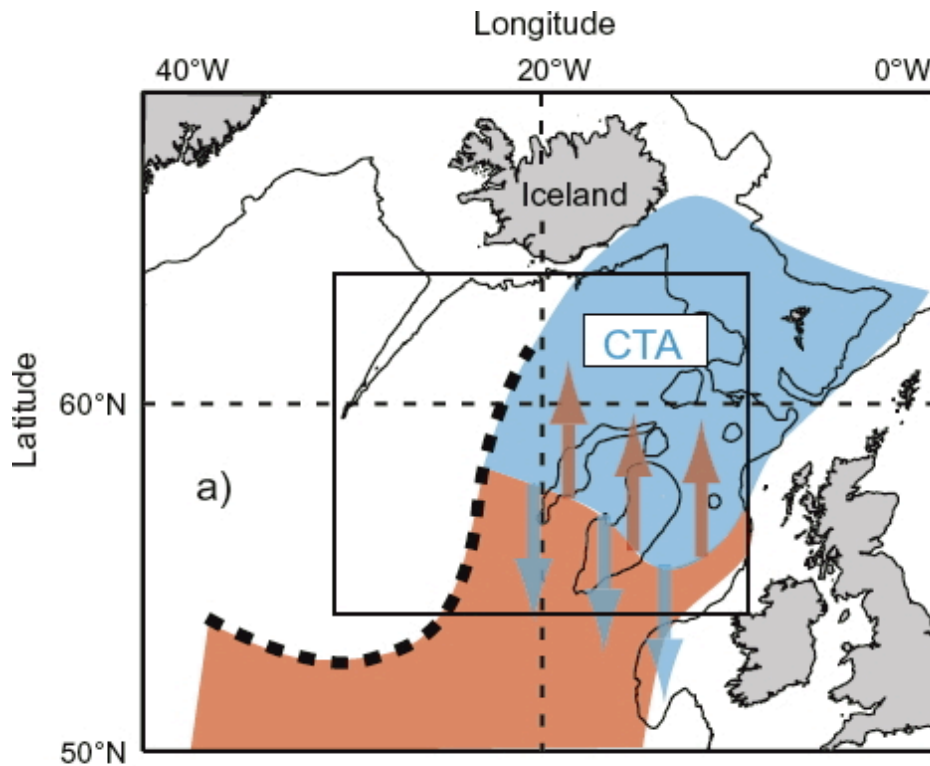


Holland et al. (2008)

...contributing to the
(dynamic) thinning of the
outlet glaciers on
Greenland
(m per year)



...and direct impact on the marine ecosystem, here abundance of zooplankton



Hatun et al, Prog. Oceanogr. (2009)

(CTA = Cold Temperature Assemblages)



Q1: How does the SPG respond to a persistent, decadal time scale positive – or negative – NAO forcing?

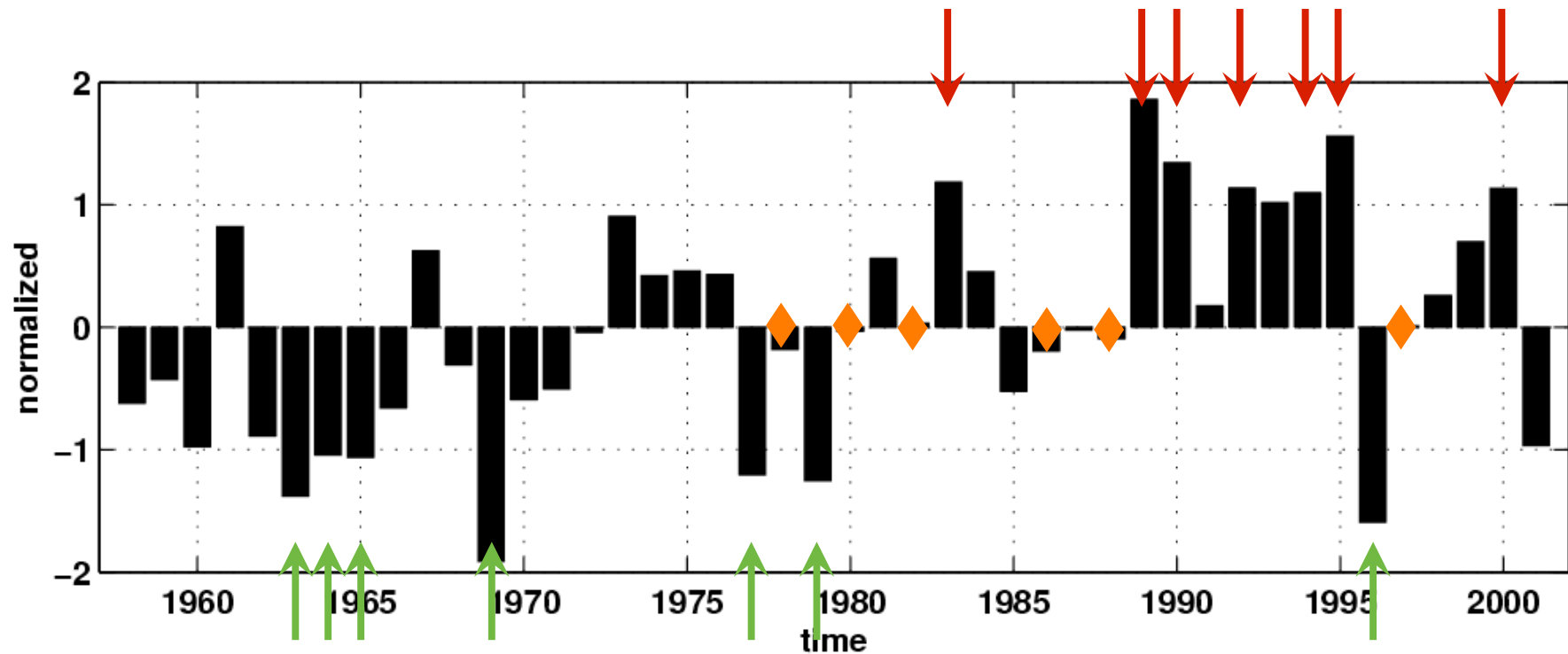
Q2: How linear is the response of the SPG forced with positive – or negative – phases of the NAO?

NAO index based on ERA40

NAO+

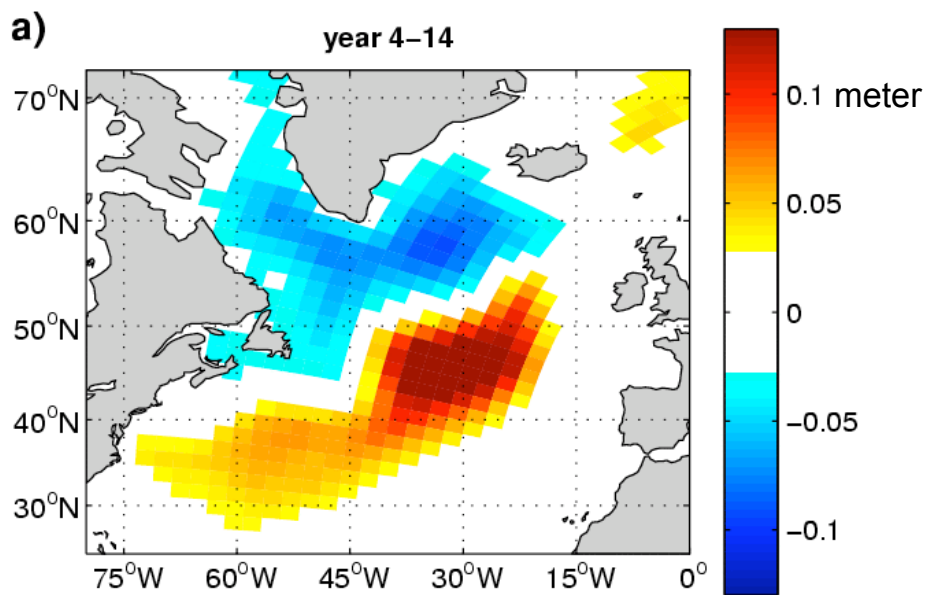
NAO-

NAOn



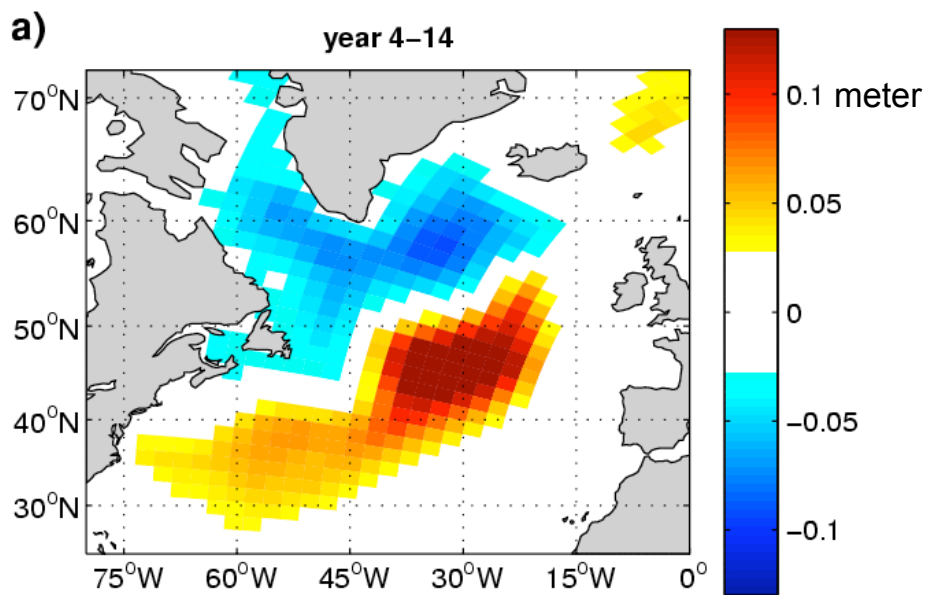
Force ocean model with repeating
(i) **NAO+**, (ii) **NAO-** or (iii) **NAOn** fields for 40 years

Sea surface height, **NAO+** minus **NAOn** (starting from 1961)

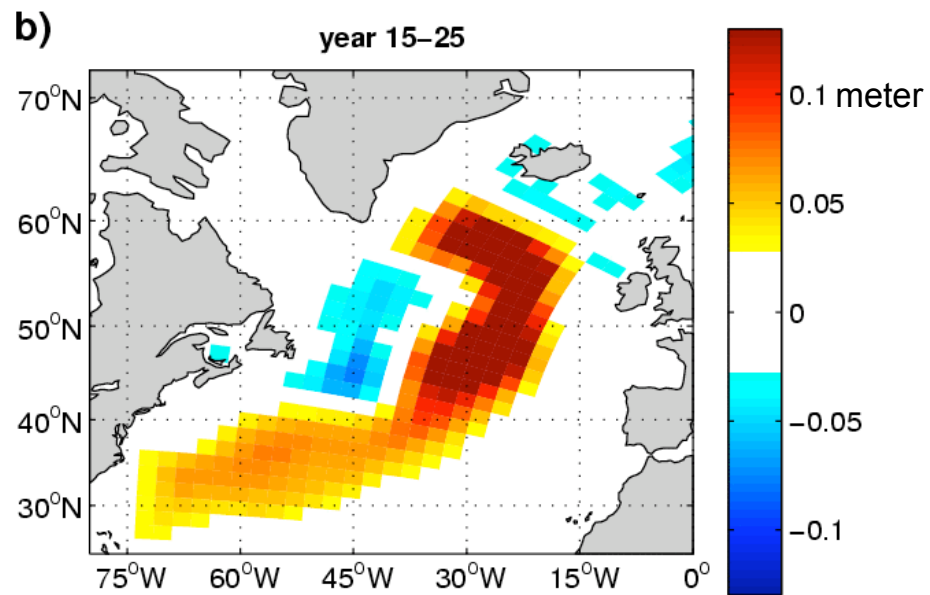


Yr 4-14: Intensified gyre

Sea surface height, **NAO+** minus **NAO_n** (starting from 1961)

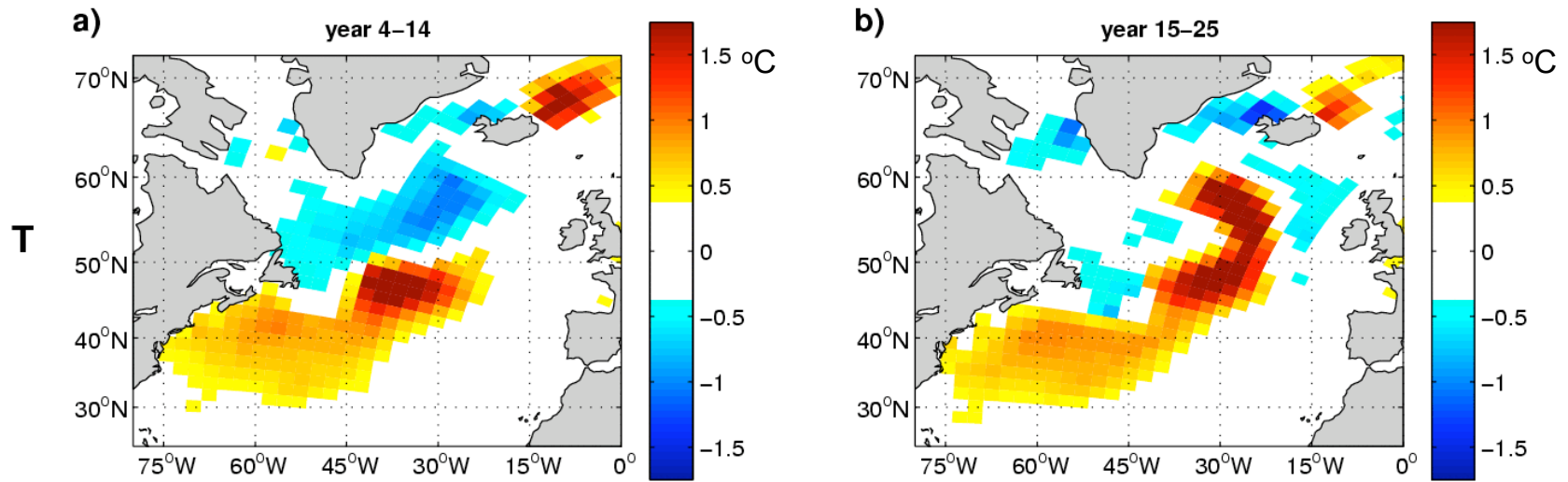


Yr 4-14: Intensified gyre

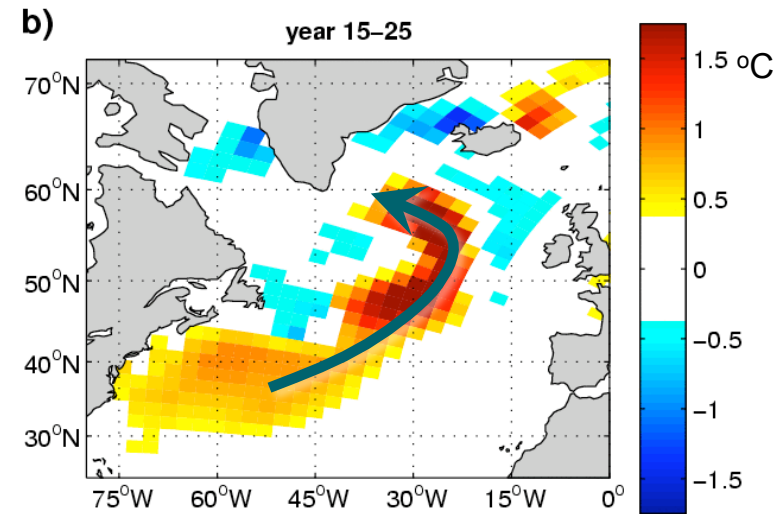
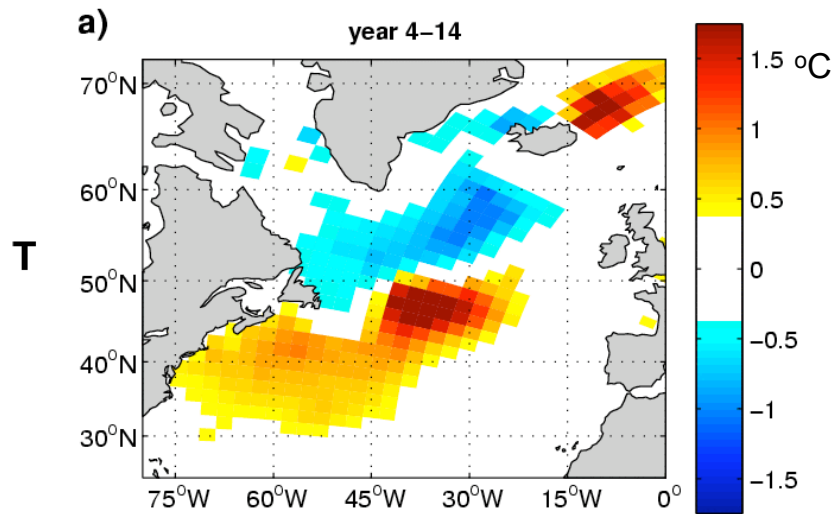


Yr 15-25: Weakened gyre

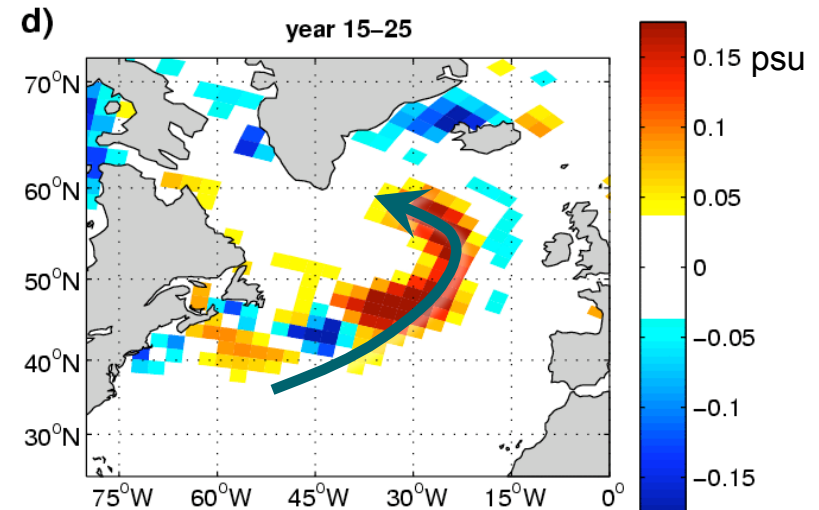
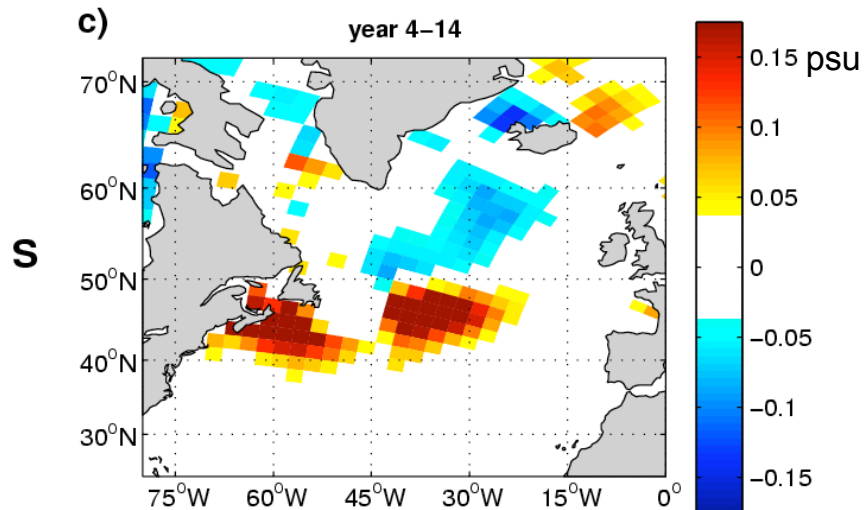
Mixed layer T and S, **NAO+** minus **NAOn** (starting from 1961)



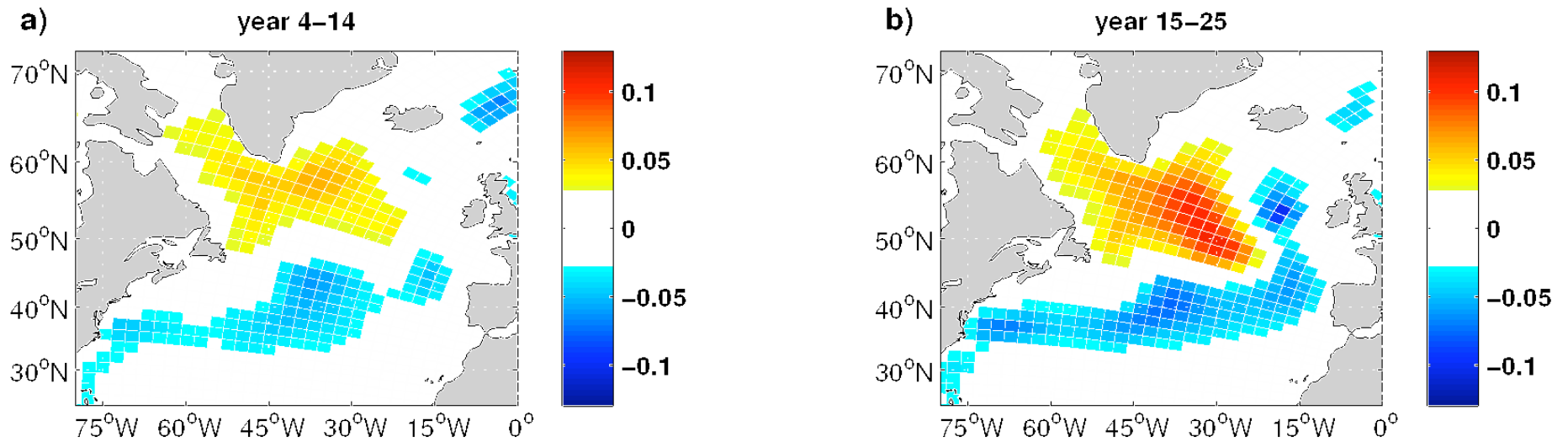
Mixed layer T and S, NAO+ minus NAO_n (starting from 1961)



Advection of warm and saline waters



Sea surface height, NAO- minus NAO_n



Gradual weakening

Conclusions (1)

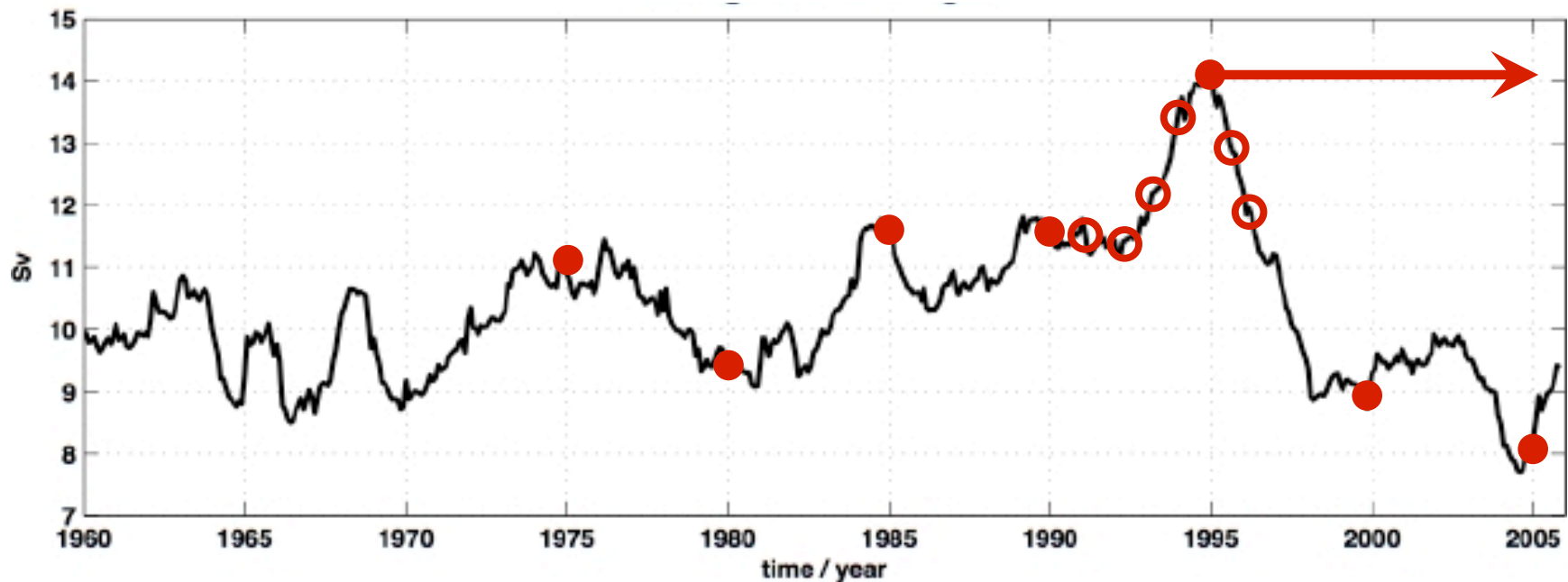
- **NAO+ forcing** →
 - Initial strengthening of SPG
 - After ~10 years replaced by weakening
 - Advection of warm water counteracts local cooling
- **NAO- forcing** →
 - Gradual weakening of SPG, approaching a minimum value
- **Asymmetry** →
 - Potentially misleading to look at (NAO+ minus NAO-)

Q3: What caused the abrupt drop in the SPG after 1995?

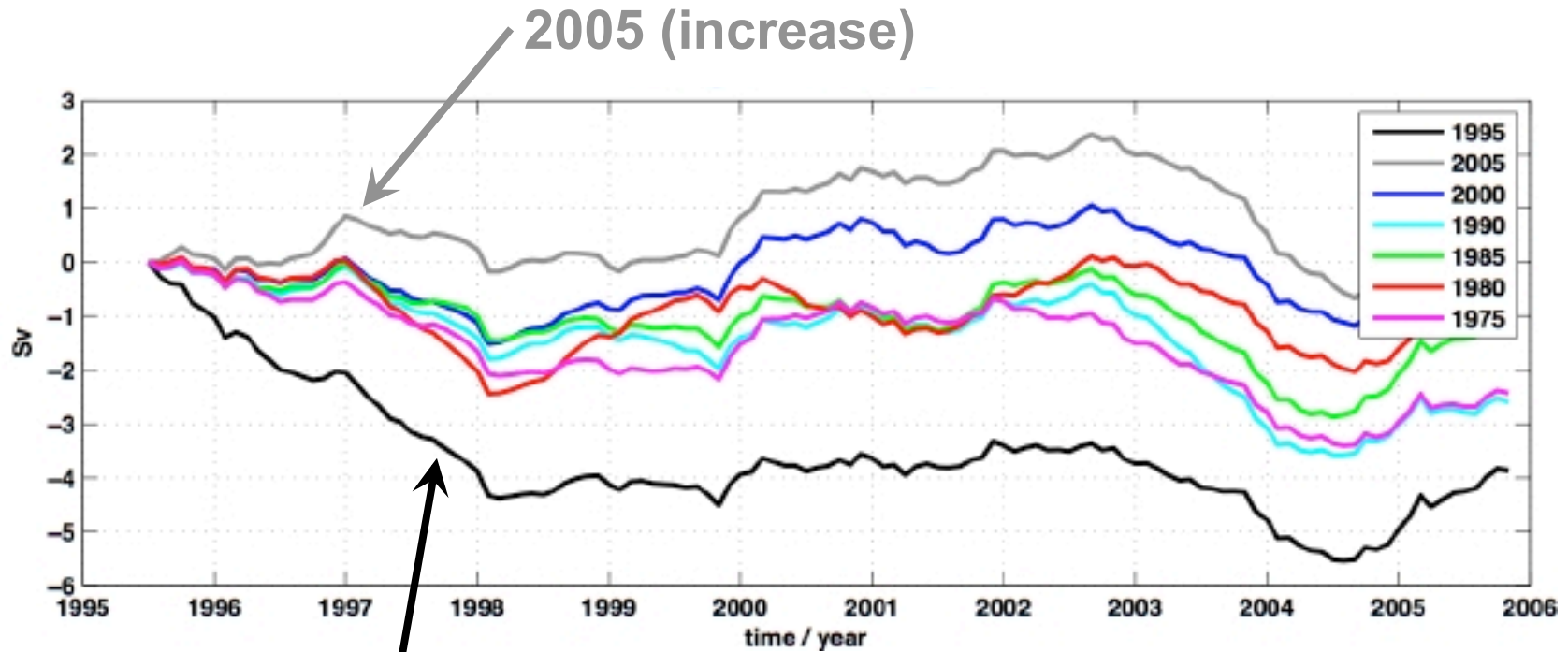
Q4: What is the role of the actual ocean initial state?

Sensitivity experiments

- same model as before.
- **post 1995 forcing** applied to initial conditions from 1975, 1980, 1985, 1990, 2000, 2005, and every year between 1991 and 1997

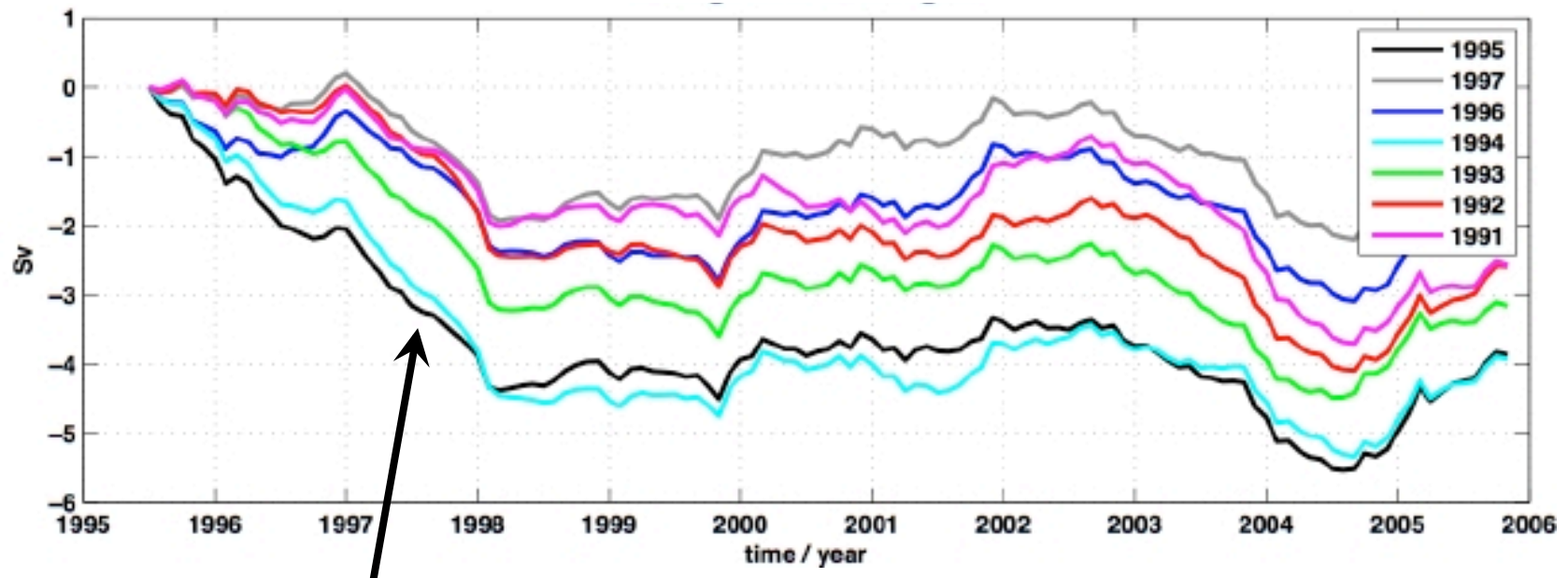


Post 1995 forcing (I)



Control integration
(1995; "real model world")

Post 1995 forcing (II)



Control integration
(1995; "real model world")

Conclusions (2)

- **SPG drop in 1995 →**

SPG at maximum strength and approaching break-down after a long period with NAO+ forcing

NAO forcing changed from high to low value the winter 1995/96

The combined effect lead to an unprecedented collapse of SPG

- **Note**

SPG would also have collapsed in 1994 with post-1995 forcing

Otherwise no collapse for the period 1960-2005 with post-95 forcing

SPG would have increased in 2005 even with post-95 forcing

- **Predictability**

Ocean initial state of crucial importance